

The age of metasedimentary rocks and their regional metamorphism in the Soursat Complex, NW IRAN: U–Pb dating of zircon and monazite, using LA–ICP–MS

Jamshidi Badr Mahboobeh ^{*,1}, Collins² Alan S, Masoudi Fariborz ^{1,3}

Corresponding author:

1. Tarbiat Moallem University, 49, Mofateh Ave, Tehran. Iran *Payame Noor University Iran *Karaj Islamic Azad University Iran.

2. Tectonics Resource and Exploration (Trax), School of Earth and Environmental Sciences, University of Adelaide, Australia

3. Shahid Beheshti University

E-mail address: m_jamshidi@tmu.ac.ir

Abstract

The Soursat Metamorphic Complex (SMC) in northwestern Iran is one of the main metamorphic terranes in northern Sanandaj-Sirjan metamorphic belt. The complex composed mainly of metamorphic rocks associated with granitic intrusions. Metamorphic rocks vary from greenschist to amphibolite facies and consist of mica-schist, garnet-schist, staurolite-schist, andalusite-schist, cordierite-schist, marble, gneiss and granite-gneiss. SMC is in tectonic contact with Precambrian to Paleozoic sedimentary rocks (Kahar, Bayandor, Soltaniyeh, Barut, Zaigon, Lalun and Mila formations) which make it difficult to date it based on stratigraphy. In this study U/Pb dating of zircons and Monazites used in order to find the ages of deposition and metamorphism of metasediments in SMC. U/Pb dating of zircons from a staurolite-schist in the complex by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) yielded a maximum depositional age of 605 ± 43 ($^{238}\text{U}/^{206}\text{Pb}$ age). Monazites were also dated from a garnet-schist using the same technique and yielded a $^{238}\text{U}/^{206}\text{Pb}$ age of 61 ± 8 which is interpreted as dating the peak of regional metamorphism. Based on these ages, metasedimentary protoliths of the Soursat Complex are interpreted to be deposited at the same time as the surrounding essentially unmetamorphosed Precambrian-Cambrian sedimentary rocks and regional metamorphism occurred later and could be related to the Paleocene orogenesis during the collision of Arabian plate with Iranian block and closure of Neotethys.

Keywords: LA-ICP-MS; Zircon; Monazite; NW Iran; Sanandaj-Sirjan; Soursat Complex.

Introduction

Precambrian terranes are exposed in many places in Iran. In the Saghand Region, Central Iran, Soursat (Soursat Metamorphic Complex), Zanjan (Takab Complex), Gorgan, Golpayegan and etc, basement complexes occur as the oldest lithostratigraphic element (e.g. Haghypour, 1974 & 1981; Nadimi, 2007; Hassanzadeh et al. 2008; Horton et al., 2008; Saki, 2009).

The term “basement complex” is used for the set of rocks underlying the Pan African unconformity and comprising mostly metamorphic or igneous rocks (with the age of the bottom of the related rock cover being variable, ranging in most cases from 570 to 500 Ma).

Soursat Metamorphic Complex (SMC) is in tectonic contact with Precambrian to Paleozoic sedimentary rocks which make it difficult to date it based on stratigraphy.

An early attempt on dating of the complex was focused on its intrusions (Pichagchi pluton) and K-Ar techniques yielded an age of 74.20 Ma (Kholghi khasraghi, 2004). This age has directly influenced palaeogeographic and plate tectonic reconstructions (Kholghi khasraghi, 1994). Recently, The U/Pb Zircon dating of intruded granitoid rocks in SCM (Jamshidi Badr et al. 2009a) presented two set of intrusions with Ediacaran-Cambrian and Palaeocene ages. Old ages from granitoids elsewhere in the same structural zone (Sanandaj-Sirjan metamorphic zone), also have been reported (Hassanzadeh et al. 2008).

In this study U-Pb Laser Inductively Coupled Mass Spectrometry (LA-ICPMS) dating of zircons and Monazites used in order to find the ages of deposition and metamorphism of metasedimentary rocks in SMC.

Geological setting

The area NW of Takab in Sanandaj-Sirjan metamorphic belt of Iran (Fig 1a) consists of sedimentary units and Soursat Metamorphic Complex. The complex is positioned above the Takab-Shahin Dezh road (Fig. 1b) and presents tectonic contact with sedimentary rocks. Two main geological units have been described for Precambrian to Paleozoic sedimentary rocks (Kholghi khasraghi, 1994): (1) upper Precambrian Kahar formation consisting of slate, sandstone and some acidic volcanic rocks those locally reveal a very low metamorphic grade, and (2) Precambrian-Cambrian and Ordovician dolomite (Bayandor and Soltaniyeh formations), sandstone, shale and dolomitic limestone (Barut, Zaigon, Lalun and Mila Formations).

SMC is mainly composed of metamorphic rocks with associated granitic intrusions. Metamorphic rocks vary from greenschist to amphibolite facies and consist of mica-schist, garnet-schist, staurolite-schist, andalusite-schist, marble, gneiss and granite-gneiss Haghypour, 1974 & 1981. Geological aspects of the complex described by many authors (e.g. Alavi-Naini et al, 1982; Kholghi khasraghi 1994; Jamshidi, 2001; Ghasemi, 2001; Kholghi Khasraghi, 2004; masoudi et al, 2006; Modjarad 2007; Ghasemi et al, 2009; Jamshidi et al, 2009).

Result and discussion

ICP-MS dating of Zircon and monazite crystals in two schists were obtained based on following basics. Zircons were separated using conventional methods that include crushing, sieving, magnetic separation and floatation. More than fifty zircon grains were handpicked under a binocular microscope. The zircons were then set in synthetic resin mounts, polished and cleaned in a warm HNO₃ ultrasonic bath. Cathodoluminescence (CL) and back-scattered electron (BSE) imaging were carried out to help characterize any compositional variation within individual zircons. Equipment and operating conditions for zircon analysis were identical to those reported by Payne et al. (2006). A spot size of 30 μm and repetition rate of 5 Hz was used for U-Pb data acquisition, producing a laser power density of ~8 J/cm². Zircon ages were calculated using the GEMOC GJ-1 zircon standard to correct for U-Pb fractionation (TIMS normalization data ²⁰⁷Pb/²⁰⁶Pb=608.3 Ma, ²⁰⁶Pb/²³⁸U=600.7 Ma and ²⁰⁷Pb/²³⁵U=602.2 Ma — Jackson et al. (2004)), and the GLITTER software for data reduction

(Van Achterbergh et al., 2001). Over the duration of this study the reported average normalized ages for GJ-1 were 609 ± 10 , 600.2 ± 2.7 and 601.9 ± 2.4 Ma for the $^{207}\text{Pb}/^{206}\text{Pb}$, $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{235}\text{U}$ ratios, respectively (n=24).

Metamorphic monazite from sample Sh-68 was imaged using backscattered electron imagery (BSE) and conducted by LA-ICPMS at the University of Adelaide. Equipment and operating conditions for monazite analysis are identical to those reported by Payne et al. (2006, 2008). U– Pb acquisition utilised 10 μm beam diameter for monazite run at a repetition rate of 5 Hz. Monazite ages were calculated using the MADEL monazite standard to correct for U– Pb fractionation (TIMS normalisation data: $^{207}\text{Pb}/^{206}\text{Pb}=490.7$ Ma, $^{206}\text{Pb}/^{238}\text{U}=514.8$ Ma, $^{207}\text{Pb}/^{235}\text{U}=510.4$ Ma: Payne et al. 2008), and again the GLITTER software for data reduction. Over the duration of this study, the reported average normalised ages for MADEL are 493.0 ± 8.3 , 514.3 ± 2.4 and 511.2 ± 2.0 Ma for the $^{207}\text{Pb}/^{206}\text{Pb}$, $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{235}\text{U}$ ratios, respectively (n=32). Accuracy was monitored by repeat analyses of the in-house internal monazite standard (94–222/ Bruna-NW: $^{206}\text{Pb}/^{238}\text{U}=447$ Ma: Payne et al. 2008). Over the duration of this study, the reported average $^{206}\text{Pb}/^{238}\text{U}$ age for the internal standard was 446.9 ± 3.1 Ma (n=15).

U/Pb dating of zircon from a staurolite-schist by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) yielded a maximum depositional age of 605 ± 43 Ma ($^{238}\text{U}/^{206}\text{Pb}$). Based on U/Pb dating of zircon, metasedimentary protoliths of the Soursat Complex are interpreted to be deposited at the same time as the surrounding essentially unmetamorphosed Precambrian-Cambrian sedimentary rocks. One group of granitoids in the complex also has Ediacaran-Cambrian age (Jamshidi Badr et al, 2009a). The same age for deposition of SMC protoliths and emplacement of a set of granitoids and sedimentation of surrounding essentially unmetamorphosed rocks support recent interpretations of researchers working elsewhere in Iran and SE Turkey in recognizing an active margin developed along this part of the Gondwanan margin (Ramezani & Tucker, 2003; Ustaomer et al. 2009) immediately after Gondwana amalgamated (Collins & Pisarevsky, 2005).

Monazite were also dated from a garnet-schist yielded a $^{238}\text{U}/^{206}\text{Pb}$ age of Palaeocene-early Eocene that could be interpreted as the time of the peak of regional metamorphism, U-Pb zircon data from another group of granitoids also yielded Palaeocene-early Eocene crystallization ages (Jamshidi Badr et al, 2009a). Regional metamorphism occurred later and it could be related to the Paleocene orogenesis during the collision of Arabian plate with Iranian block and closure of Neotethys.

Reference

1. Alavi-Naini, M., Hajian, J., Amidi, A., Bolurchi, H., 1982. Geology of Takab-Saein Qaleh: Explanatory note of 1:250000 map of Takab quadrangle, Geological Survey of Iran, Report No.50.
2. Collins, A.S. and Pisarevsky, S.A., 2005. Amalgamating eastern Gondwana: The evolution of the Circum-Indian Orogens. *Earth Science Reviews*, v/ 71, p 229-270.
3. Ghasemi, A., 2001, Structural Study of south Shahin-Dej Metamorphic rocks of NW Iran. Ph.D.Thesis (in Farsi), Shahid beheshti university.
4. Ghasemi, A., Poor Kerman, M., 2009. Structure of the Soresat Metamorphic Complex, North Sanandaj-Sirjan Zone, northwest Iran, *Australian Journal of Earth Sciences*, 56, v. 7, p. 939 - 949.

5. Haghypour, A. 1974. Etude géologique de la région de Biabanak-Bafg (Iran Central); pétrologie et tectonique du précambrien et de sa couverture, Ph.D. thesis, université scientifique et médicale de Grenoble, France, 403.
6. Haghypour, A., 1981. Precambrian in central Iran: lithostratigraphy, structural history and petrology. Iranian Petroleum Institute Bulletin, 81, p. 1–17.
7. Hassanzadeh, J., Stockli, D. F., Horton, B. K., Axen, G. J., Stockli, L. D., Grove, M., Schmitt, A. K., Walker, J. D., 2008. U-Pb zircon geochronology of late Neoproterozoic–Early Cambrian granitoids in Iran: Implications for paleogeography, magmatism, and exhumation history of Iranian basement. Tectonophysics, 451, p. 71–96.
8. Horton, B.K., Hassanzadeh, J., Stockli, D.F., Axen, G.J., Gillis, R.J., Guest, B., Amini, A., Fakhari, M., Zamanzadeh, M.S., Grove, G., 2008. Detrital zircon provenance of Neoproterozoic to Cenozoic deposits in Iran: Implications for chronostratigraphy and collisional tectonics. Tectonophysics 451, p. 97–122.
9. Jamshidi badr, M., 2001. Petrology and Petrography study of metamorphic and Igneous rocks of Shahindezh area, M.S.C. Thesis (in Farsi), Tabriz university of Iran, 140p.
10. Jamshidi badr, M., Collins, A., Masoudi. F., Mehrabi, B., 2009a. detrital u-pb zircon dating of granitoids in the soursat complex (NW IRAN). Geological Society of America Abstracts with Programs, v41, N. 7, p. 113.
11. Jamshidi badr, M., Masoudi. F., Mohajjel, M., 2009b. State and condition of the formation of cordierite crystal in Metapelites of Soursat Complex, the 17th symposium of society of Crystallography and Mineralogy of Iran.
12. Masoudi. F., Mohajjel, M., Jamshidi badr, M., 2006. Study of different schistosity in Soursat complex based on distribution of metamorphic minerals, 14th Symposium of Crystallography and Mineralogy of Iran.
13. Modjarad, M, 2007. Petrology and geodynamic rocks of sursat (east of shahindezh) NW Iran, Ph.D.Thesis (in Farsi), Tabriz University.
14. Nadimi, A., 2007. Evolution of the central Iranian basement. Gondwana Research 12, v 3, p 324–333.
15. Payne, J. L., Hand, M., Barovich, K. M. & Wade, B. P., 2008. Temporal constraints on the timing of high-grade metamorphism in the northern Gawler Craton: implications for assembly of the Australian Proterozoic. Australian Journal of Earth Sciences 55, p. 623–640.
16. Payne, J.L., Barovich, K., Hand, M., 2006. Provenance of metasedimentary rocks in the northern Gawler Craton, Australia: implications for Palaeoproterozoic reconstructions. Precambrian Research, 148, 275–291.
17. Ramezani, J. and Tucker, R.D., 2003. The Saghand region, central Iran: U-Pb geochronology, petrogenesis and implications for Gondwana tectonics. American Journal Of Science, 303, p. 622-665.
18. Saki, A., 2009. Proto-Tethyan remnants in northwest Iran: Geochemistry of the gneisses and metapelitic rocks. Gondwana Research. doi:10.1016/j.gr.2009.08.008.
19. Ustaömer, P.A., Ustaömer, T., Collins, A.S. and Robertson, A.H.F., 2009. Cadomian (Ediacaran–Cambrian) arc magmatism in the Bitlis Massif, SE Turkey: Magmatism along the developing northern margin of Gondwana. Tectonophysics, 473, p. 99–112.

